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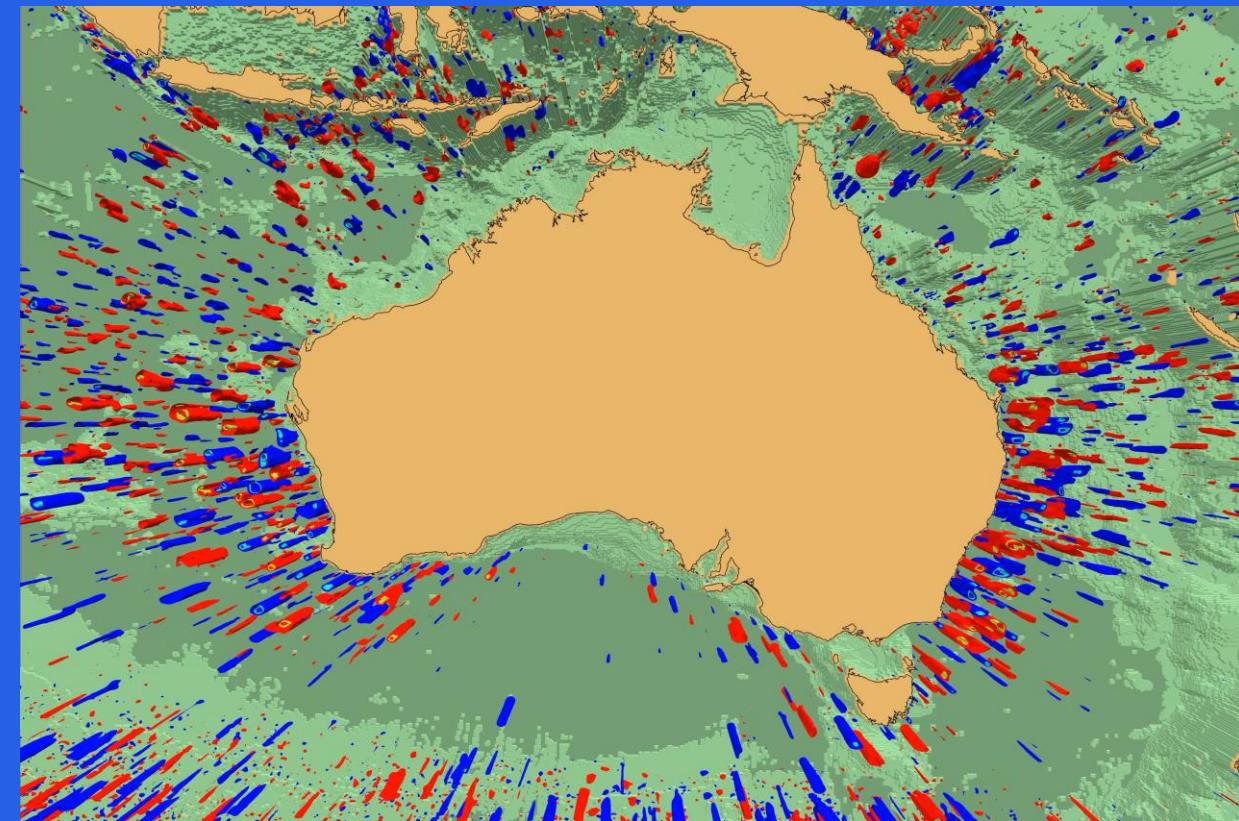
An operational ensemble Kalman Filter global ocean forecasting system OceanMAPSv4

Gary Brassington

Pavel Sakov, Prasanth Divakaran, Saima Aijaz, Mirko Velic,
Chris Bladwell, Matt Chamberlain, Xinmei Huang, Roy
Thompson, Mikhail Entel, Ahmad Khan



Bluelink
Ocean Forecasting





Overview

Bluelink since 2002 – 20+ years of ocean forecasting science

Why Bluelink? Defence and National benefit (Commercial/Public/Research)

In 2019 we faced a design choice

Higher model resolution and 3D DA (preparation for SWOT)

Persist model resolution and implement 4D DA

4D DA – Asynchronous Hybrid-DLETKF

- EnKF vs EnOI benefits
- OceanMAPSv4 performance

Conclusion





OceanMAPSv4.1i

System

Model

OFAM3 (MOM5)
75S-75N, 0-360
0.1° x 0.1°, 51 z*-levels (5m top cell)

Data assimilation

EnKF-C / Hybrid-DLETKF
48 dynamic members
144 low-mode members
FGAT, Restart initialisation

Atmospheric forcing

ACCESS-G4 (APS4)
Bulk formulae

Observations

In situ profiles (GTS, GDAC)
Satellite altimetry (RADS, J3, SARAL,
Sentinel-3A, Cryosat-2)
Satellite SST (AMSR2, NAVOCEANO,
NPP-VIIRS, NOAA20-VIIRS)

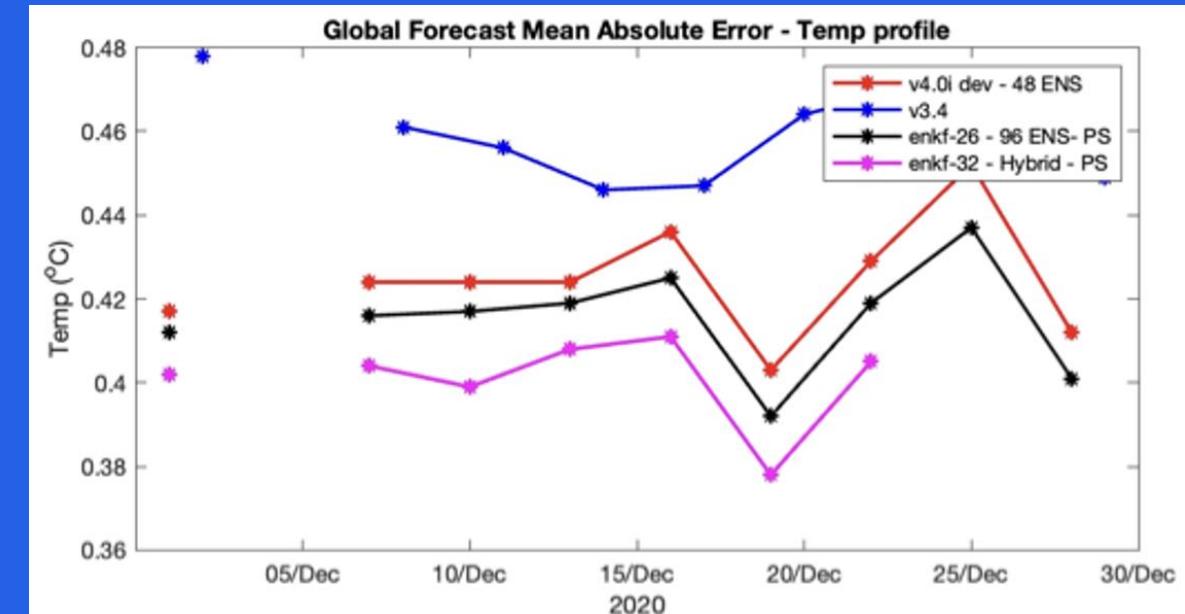
Forecasts

EnKF BRT (-3 day analysis)
EnKF NRT (-2 day analysis)
Daily 7 day forecasts (001, Av. RES)

Impact

Performance

- Reduced forecast innovation
- Reduced latency (-2 days)
- More dynamically balanced increments
- Uniform latency
- Beats persistence
- Reduction in abyssal noise
- Improved separation of eddies
- Sustaining low signal to noise eddies
- Improved skill ocean currents



References

- Brassington, G. B., P. Sakov, P. Divakaran, S. Aijaz, J. Sweeney-Van Kinderen, X. Huang and S. Allen, 2023, June. OceanMAPS v4. 0i: a global eddy resolving EnKF ocean forecasting system. In *OCEANS 2023-Limerick* (pp. 1-8). IEEE
Sakov, P., 2014. EnKF-C user guide. *arXiv preprint arXiv:1410.1233*.



The forecast stage of the EnKF involves just propagating each ensemble member:

$$\mathbf{E}_i^f = \mathcal{M}_i(\mathbf{E}_{i-1}^a). \quad (1.13)$$

$$\mathbf{X} = \{\mathbf{E}\} = \{\mathbf{x}, \mathbf{A}\}. \quad (1.9)$$

It is related to the SDAS of the KF (1.8) as follows:

$$\mathbf{x} = \frac{1}{m} \mathbf{E} \mathbf{1}, \quad (1.10a)$$

$$\mathbf{P} = \frac{1}{m-1} \mathbf{A} \mathbf{A}^T, \quad (1.10b)$$

$$\mathbf{A} \equiv \mathbf{E} - \mathbf{x} \mathbf{1}^T, \quad (1.10c)$$

Analysis update on the ensemble mean

$$\mathbf{x}_{k+1}^a = \mathbf{x}_{k+1}^f + \mathbf{K}_{k+1} \left[\mathbf{y}_{k+1} - \mathcal{H}_{k+1}(\mathbf{x}_{k+1}^f) \right], \quad (1.6a)$$

where

$$\mathbf{K}_{k+1} \equiv \mathbf{P}_{k+1}^f (\mathbf{H}_{k+1})^T \left[\mathbf{H}_{k+1} \mathbf{P}_{k+1}^f (\mathbf{H}_{k+1})^T + \mathbf{R}_{k+1} \right]^{-1} \quad (1.6c)$$

or $\mathbf{K} = \mathbf{A}(\mathbf{H}\mathbf{A})^T [\mathbf{H}\mathbf{A}(\mathbf{H}\mathbf{A})^T + \mathbf{R}]^{-1}$
equivalently

Ensemble anomaly update (model error covariance)

$$\mathbf{A}^a = \mathbf{T}_L \mathbf{A}^f, \quad (1.15)$$

where

$$\mathbf{T}_L = \mathbf{I} - \frac{1}{2} \mathbf{K} \mathbf{H}. \quad (1.30)$$

$$\mathbf{E}^a = \mathbf{x}^a \mathbf{1}^T + \mathbf{A}^a \quad (1.11a)$$

EnKF-C

- Hunt, B. R., E. J. Kostelich and I. Szunyogh, 2007: Efficient data assimilation for spatiotemporal chaos: A local ensemble transform filter. *Physica D*, 230, 112-126.
 Sakov, P. and P. R. Oke, 2008: A deterministic formulation of the ensemble Kalman filter: an alternative to ensemble square root filters, *Tellus*, 60A, 361-371
 Sakov, P., 2014. EnKF-C user guide. *arXiv preprint arXiv:1410.1233*.

Forecast error covariance

$$\mathbf{P}^f = \frac{1}{n-1} \mathbf{A} \mathbf{A}^T$$

where $\mathbf{A} = \partial [a_1, a_2, \dots, a_n]$

$a_i = \bar{T}_i^{month} - \bar{T}_i^{clim}$	Stationary
$a_i(t) = T_i(t) - \bar{T}(t)$	Dynamic

Observation error covariance

$$\mathbf{R} = S_{obs}^2 \mathbf{I} + S_{rep}^2 \mathbf{I}$$

Real-time observations

- Satellite altimetry (RADS)
- Satellite SST (VIIRS, AMSR2)
- In situ profiles (Argo, CTD, XBT)

Oke, P.R. and Sakov, P., 2008. Representation error of oceanic observations for data assimilation. Journal of Atmospheric and Oceanic Technology, 25(6), pp.1004-1017.

Hybrid implementation

$$\mathbf{P}^f = \partial \mathbf{P}^{dyn} + g \mathbf{P}^{stat}$$

$$\partial = 1.03 \quad g = 0.25$$

$$\mathbf{P}^{dyn} = A^{dyn} (A^{dyn})^T / (m^{dyn} - 1)$$

$$m^{dyn} = 48$$

$$\mathbf{P}^{stat} = A^{stat} (A^{stat})^T / (m^{stat} - 1)$$

$$m^{stat} = 144$$

Chamberlain, M. A., P. R. Oke, G. B. Brassington, P. Sandery, P. Divakaran, R. A. S. and Fiedler, 2021: Multiscale data assimilation in the Bluelink ocean reanalysis (BRAN). Ocean Modelling, 166, p.101849.

Asynchronous data assimilation

$$[\text{SLA}, \text{SST}, \text{T}, \text{S}] = [12\text{h}, 6\text{h}, 24\text{h}, 24\text{h}]$$

averaging bin sizes exploit satellite/2D

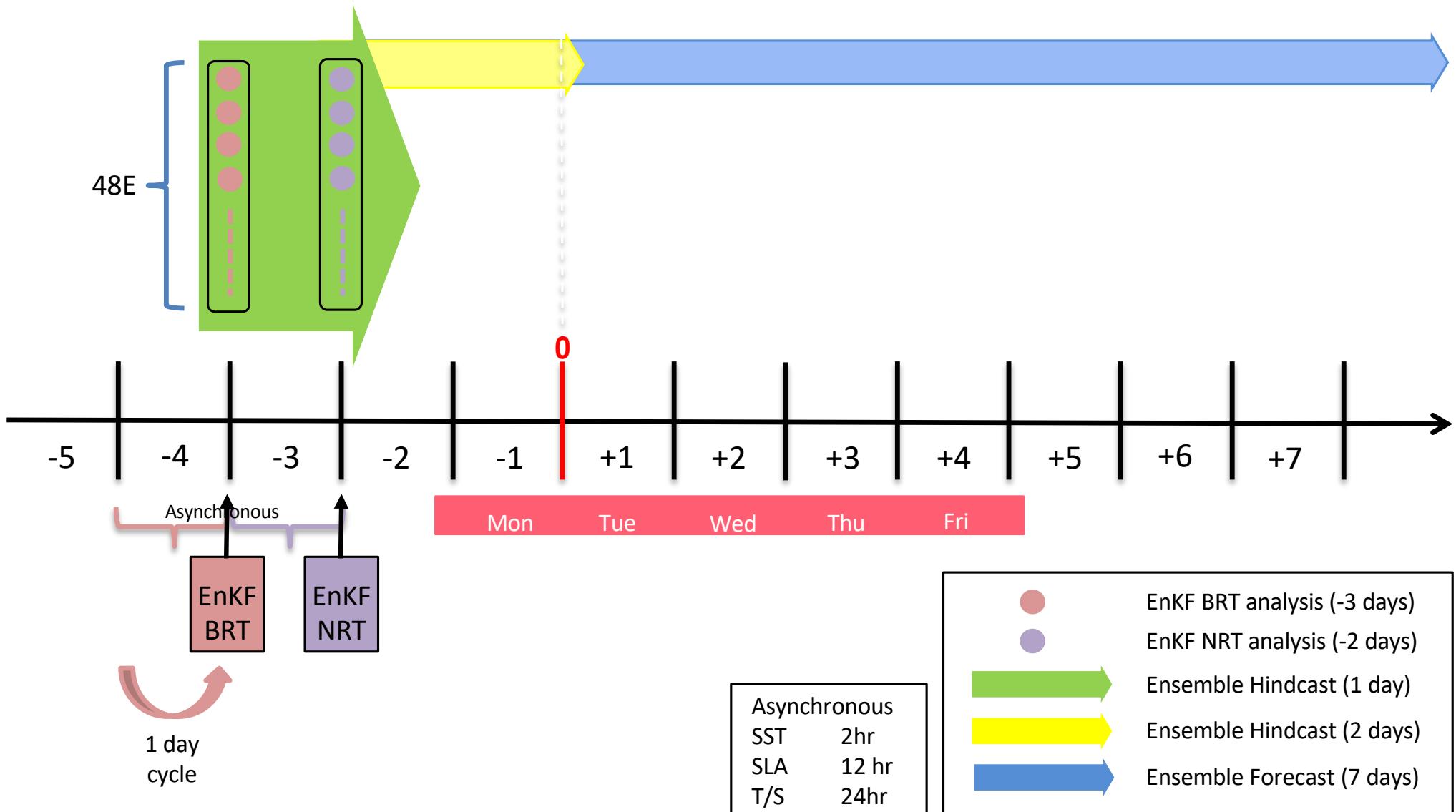
Sakov, P., Evensen, G. and Bertino, L., 2010. Asynchronous data assimilation with the EnKF. Tellus A: Dynamic Meteorology and Oceanography, 62(1), pp.24-29.

Localisation

$$[\text{SLA}, \text{SST}, \text{T}, \text{S}] = [175, 150, 450, 450] \text{ km}$$

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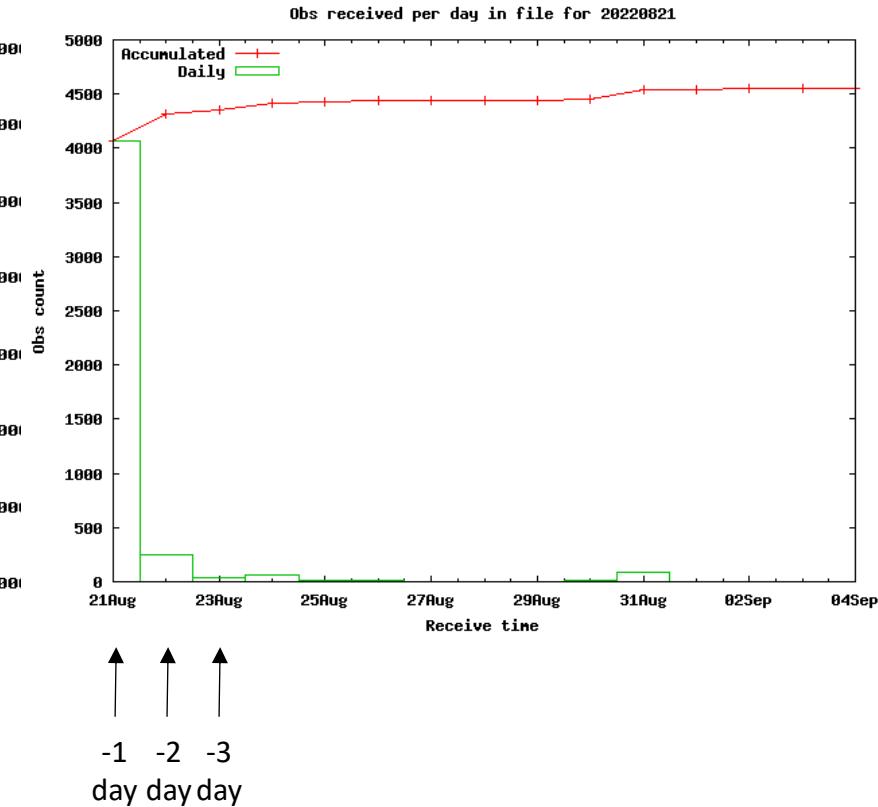
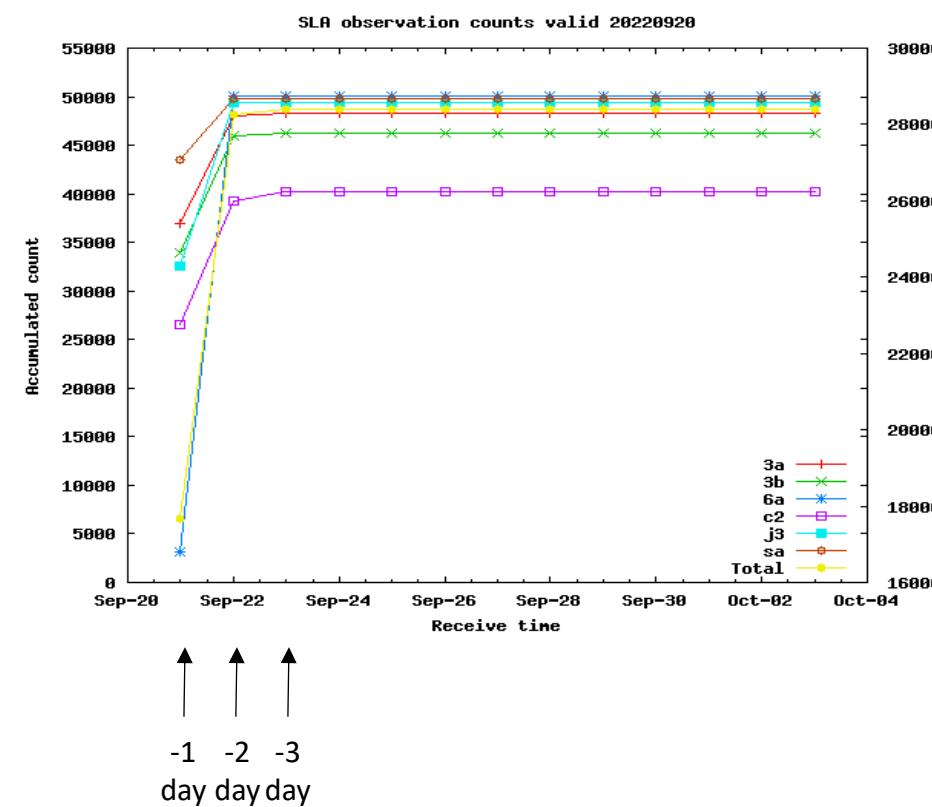
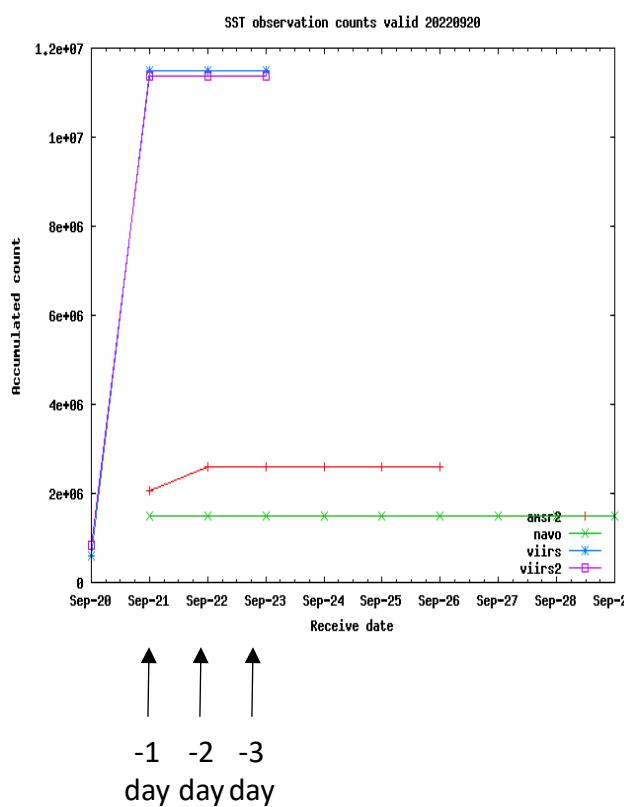
MEM001
Av. restart



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Realtime observation latency



Satellite SST
Satellite altimetry
Profiles

Low latency/Good coverage
Medium latency/Medium coverage
Medium latency/Sparse coverage

-3 days – persistent analysis cycle
-2 days – intermediate analysis to introduce available obs
-1 days – imbalanced obs not tested





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EnKF vs EnOI

Geostrophic increments

$$u_{inc}^g(z) = \frac{g}{f\rho_0} \frac{\partial}{\partial y} \int_{-H}^z \Delta\rho_{inc} dz$$

$$v_{inc}^g(z) = -\frac{g}{f\rho_0} \frac{\partial}{\partial x} \int_{-H}^z \Delta\rho_{inc} dz$$

How consistent are the density increments with geostrophic increments?

How much ageostrophic increment remains?

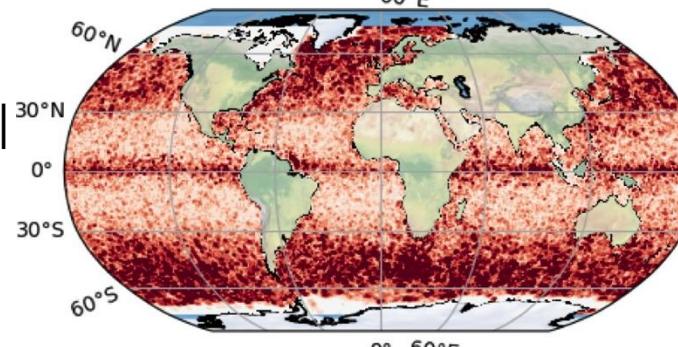
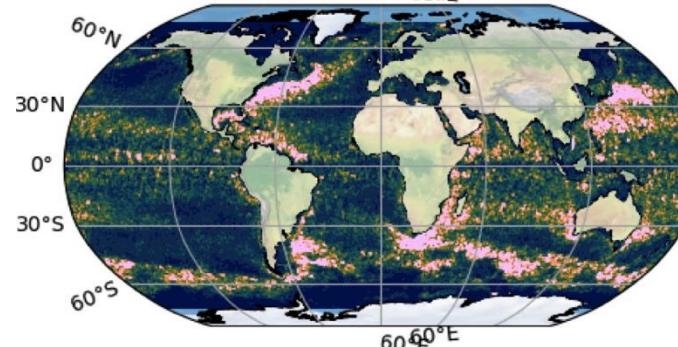
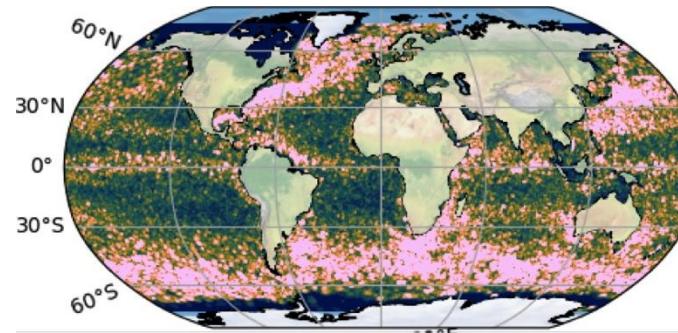
* $\Delta\rho_{inc} = \rho(T + T_{inc}, S + S_{inc}) - \rho(T, S)$,
-H is the depth of the water column

$|u_{inc}|$

$|u_{inc}^g|$

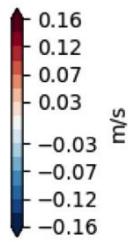
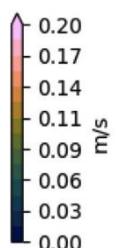
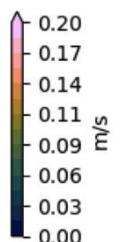
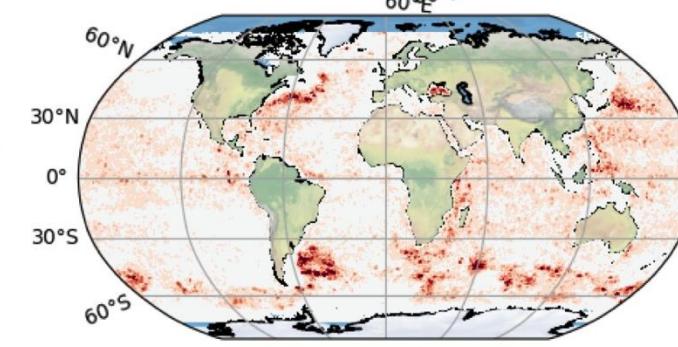
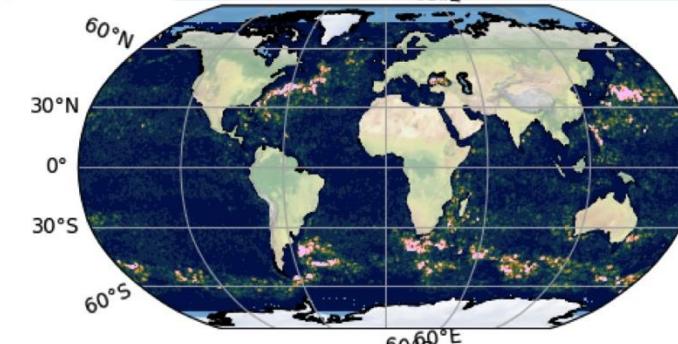
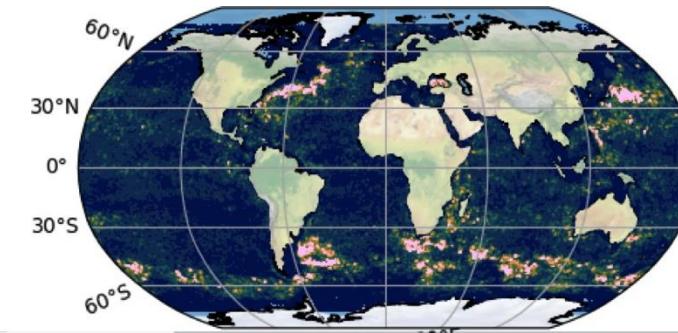
$|u_{inc}| - |u_{inc}^g|$

V3.4 EnOI



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EnKF MBR 001



Pseudo steric sea level

$$\eta_{inc}^{steric} = -\frac{1}{\rho_0} \int_{-H}^{\eta} \Delta\rho_{inc} dz$$

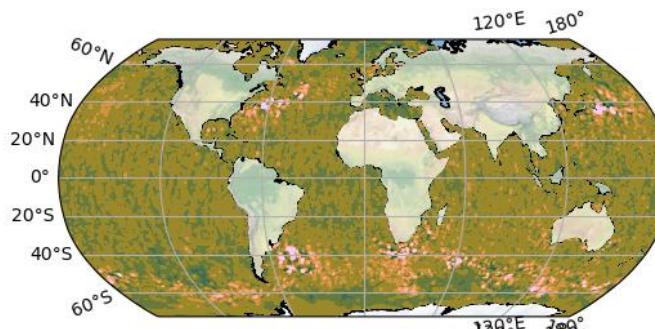
$$\Delta\eta_{inc} = \eta_{inc} - \eta_{inc}^{steric}$$

How consistent are the density increments with pseudo steric sea level?

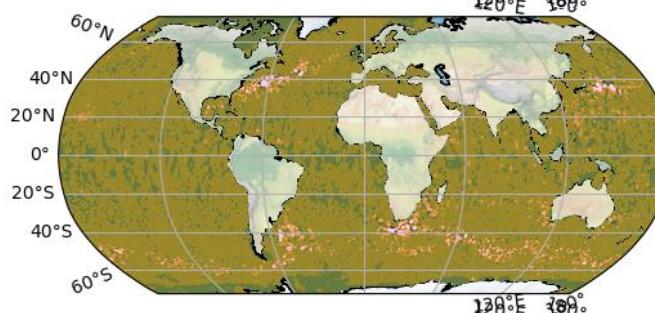
How much sea level increment contributed to the shallow water equation?

V3.4 EnOI

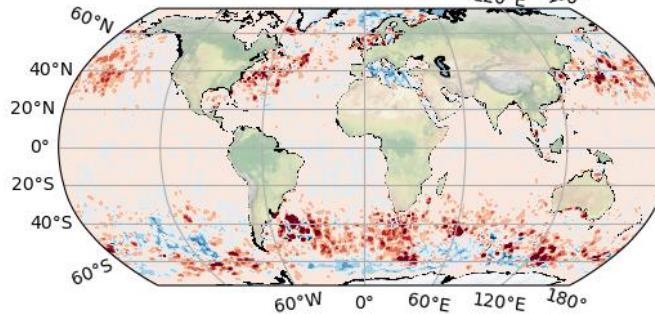
Total η_{inc}



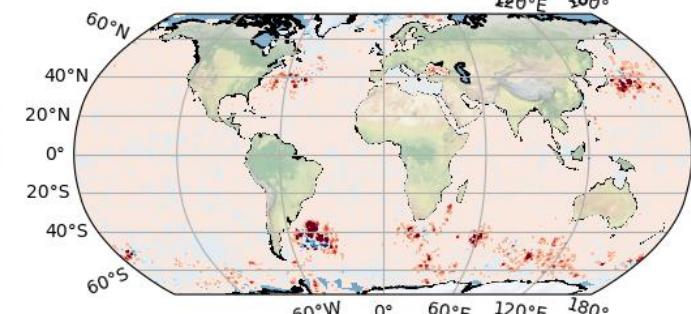
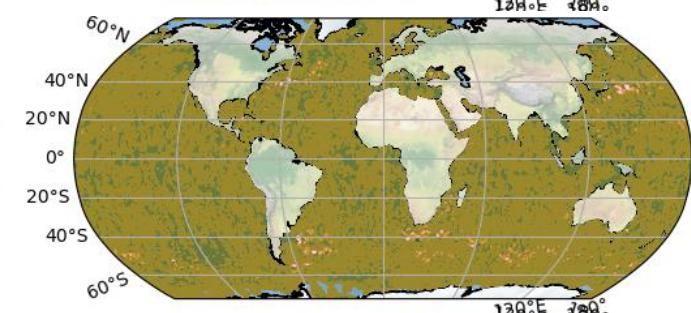
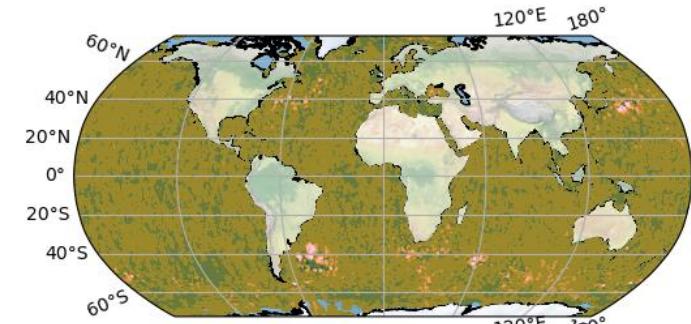
Pseudo
steric
 η_{inc}^{steric}



Difference
 $\Delta\eta_{inc}$



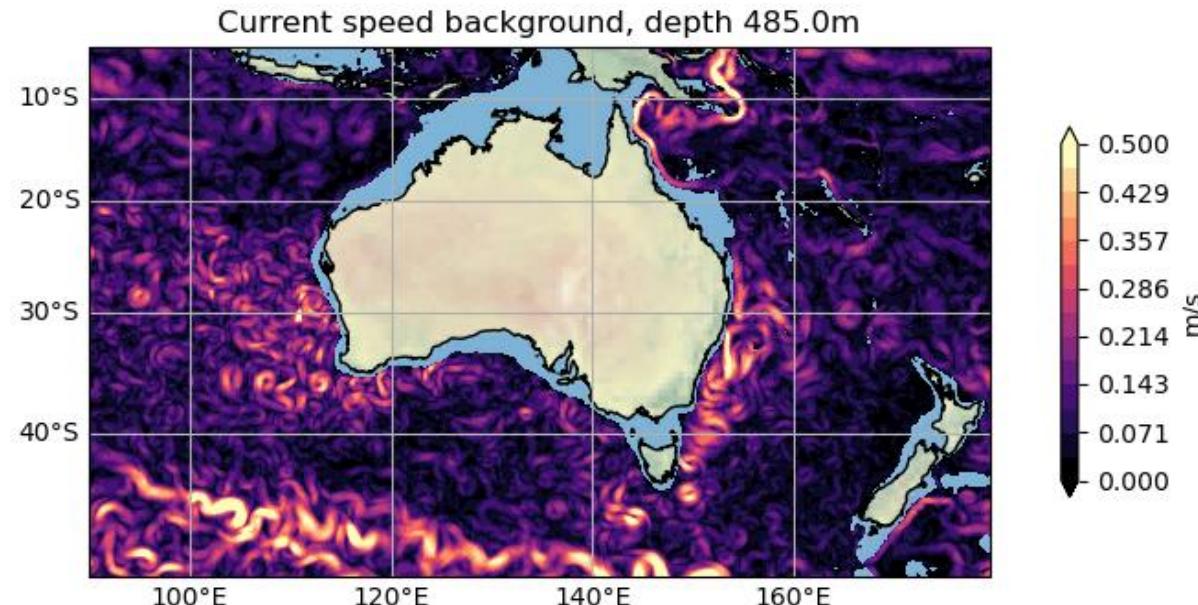
EnKF MBR 001



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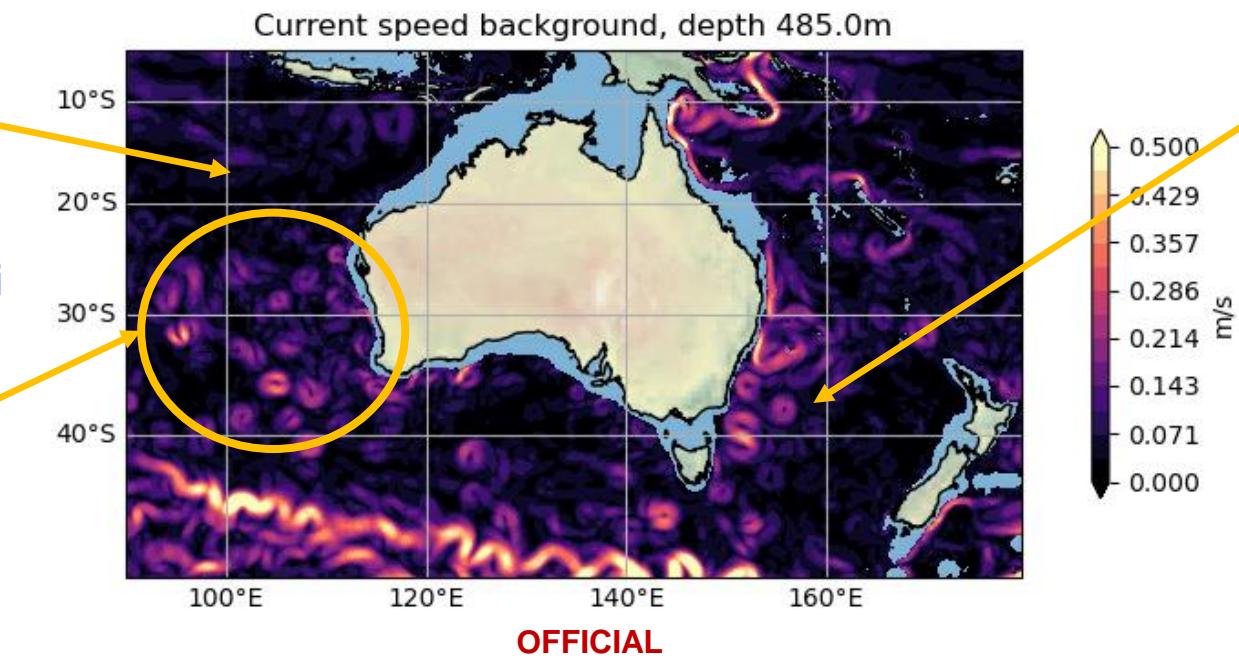
OceanMAPSv3.4
(EnOI)



Realistic
passive
regions

OceanMAPSv4.0i
(EnKF)

Clearly defined
eddies that are
observed



Clear
separation of
eddies

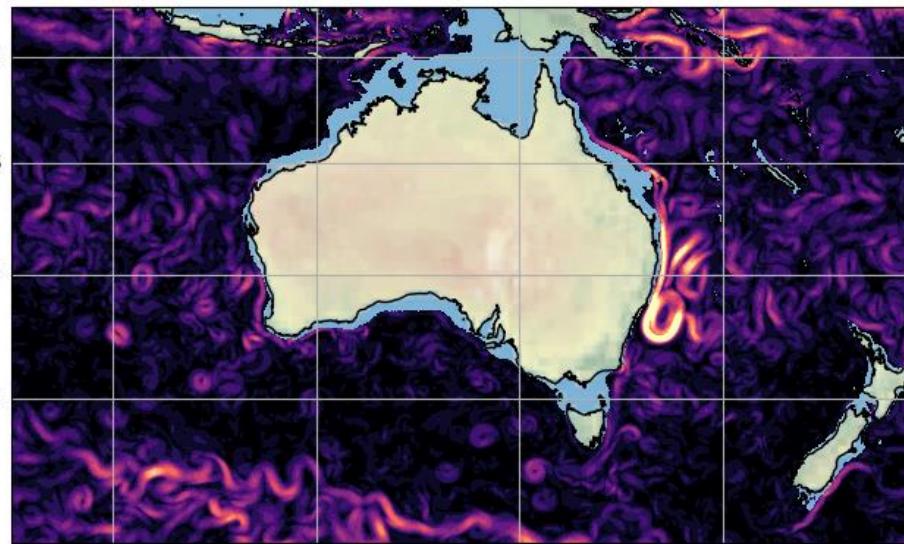
3-day hindcast
Ocean current speed
485 m
25 JULY 2020

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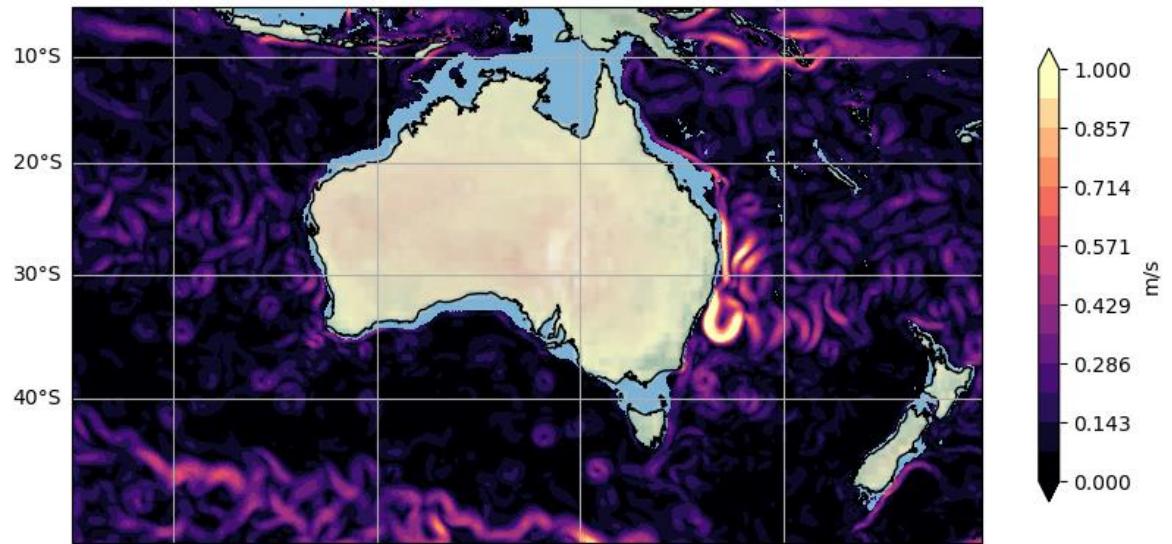


95m

enkf-MBR001 20210211 current speed background, depth 95.0m

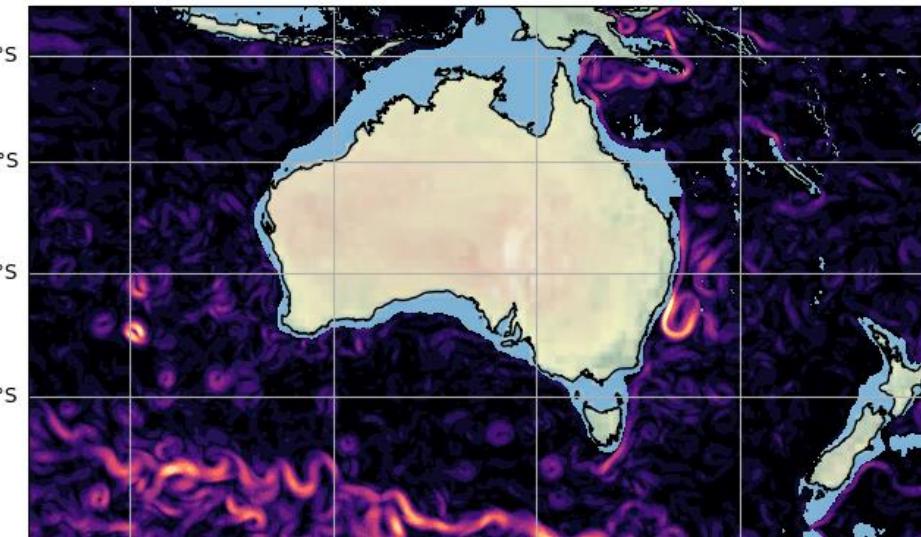


enkf-ensmean 20210211 current speed background, depth 95.0m

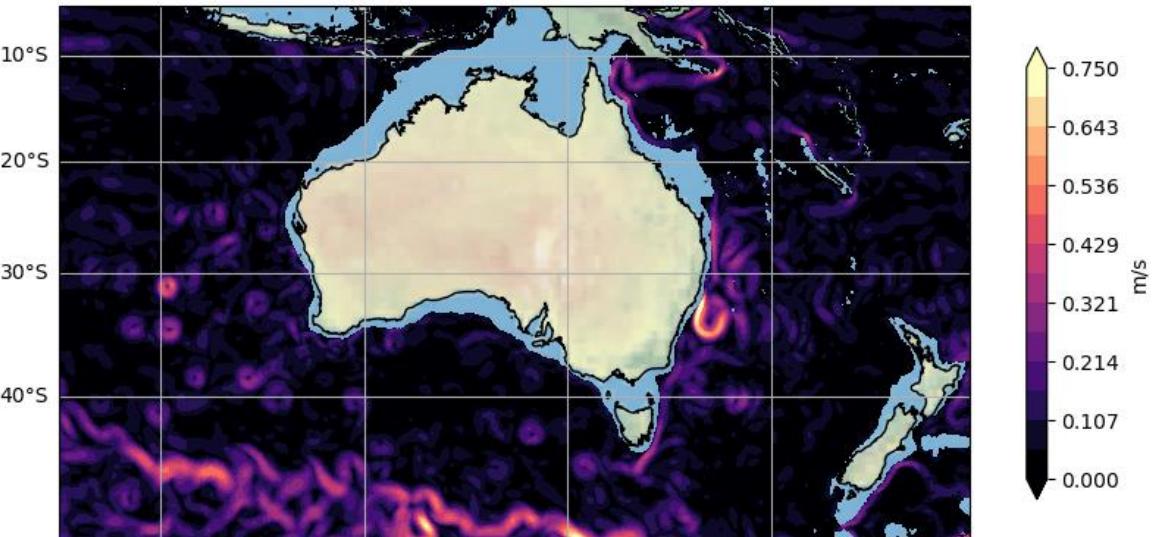


485m

enkf-MBR001 20210211 current speed background, depth 485.0m



enkf-ensmean 20210211 current speed background, depth 485.0m



MEM001

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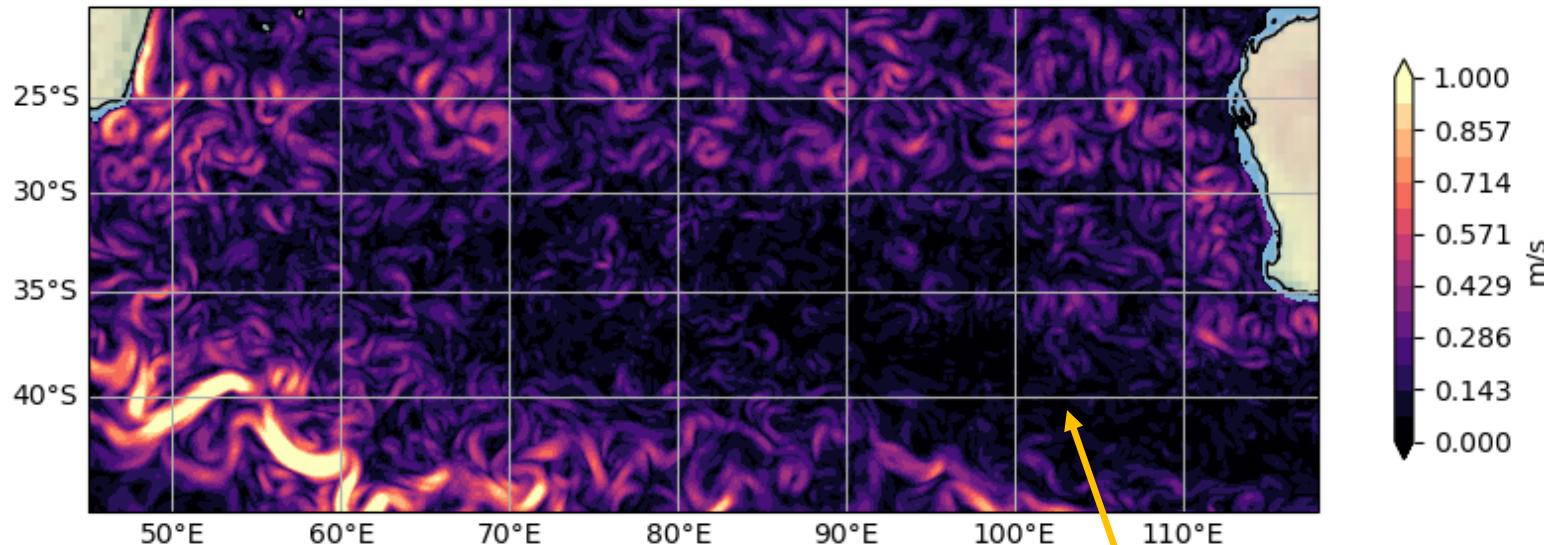
ENSMEAN

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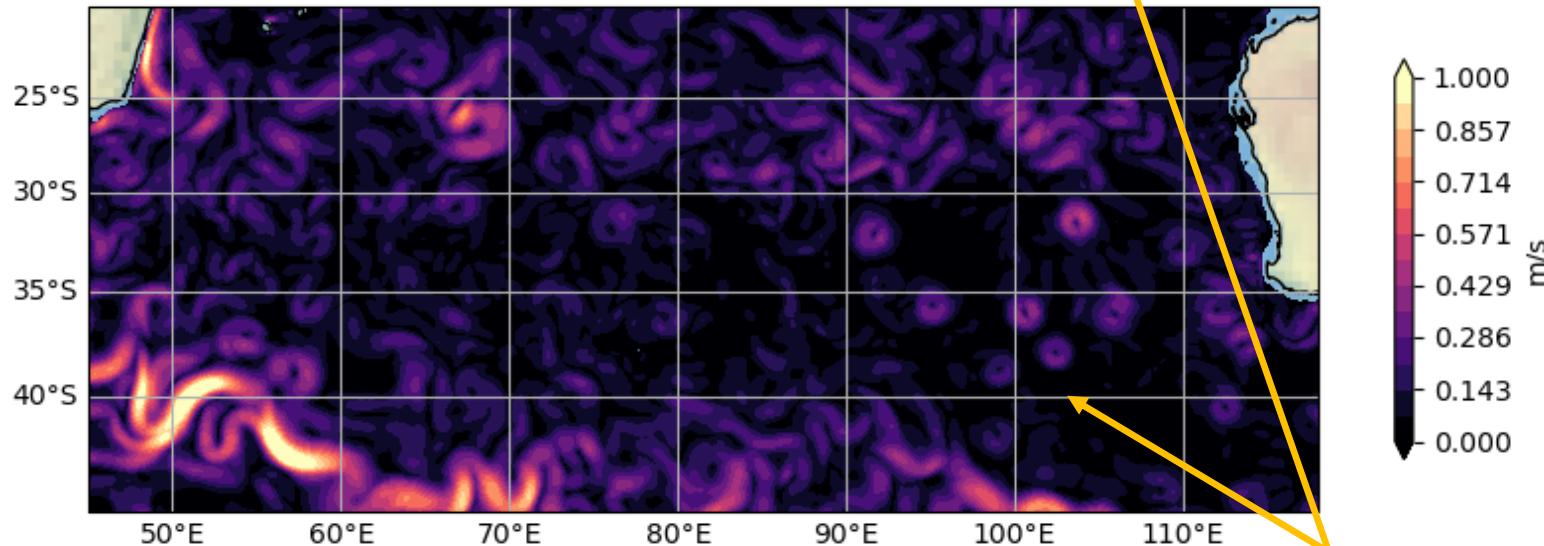
Current speed background, depth 95.0m

EnOI



Current speed background, depth 95.0m

EnKF

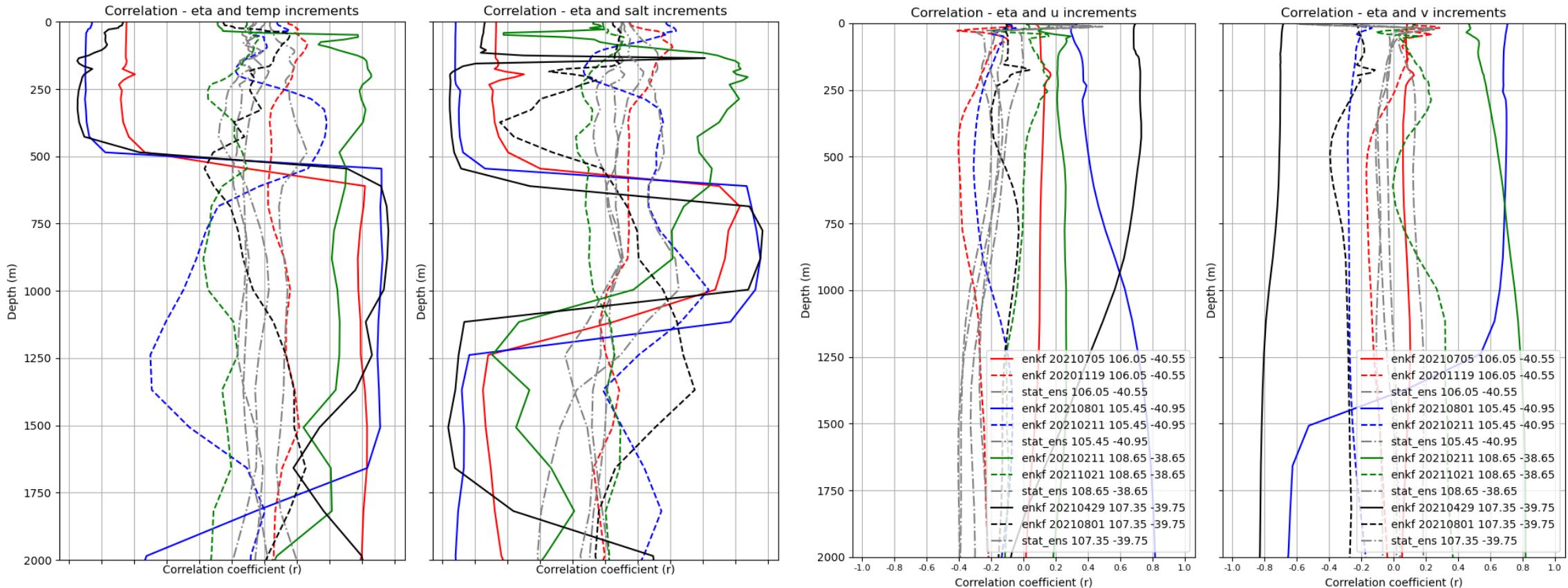


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EnOI dissipates low signal to noise eddies



CORRELATION COEFFICIENTS in the presence/absence of eddies



- Four eddies from SE Indian Ocean eddy graveyard.
- Identified through SLA signature in EnKF ensmean.
- All are cold-core eddies.

Solid lines are correlations for EnKF ensemble increments when eddy is present
Dashed lines are correlations for EnKF ensemble increments at same location on a different day (no eddy)
Grey lines are correlations for EnOI stationary ensemble members at the same location.

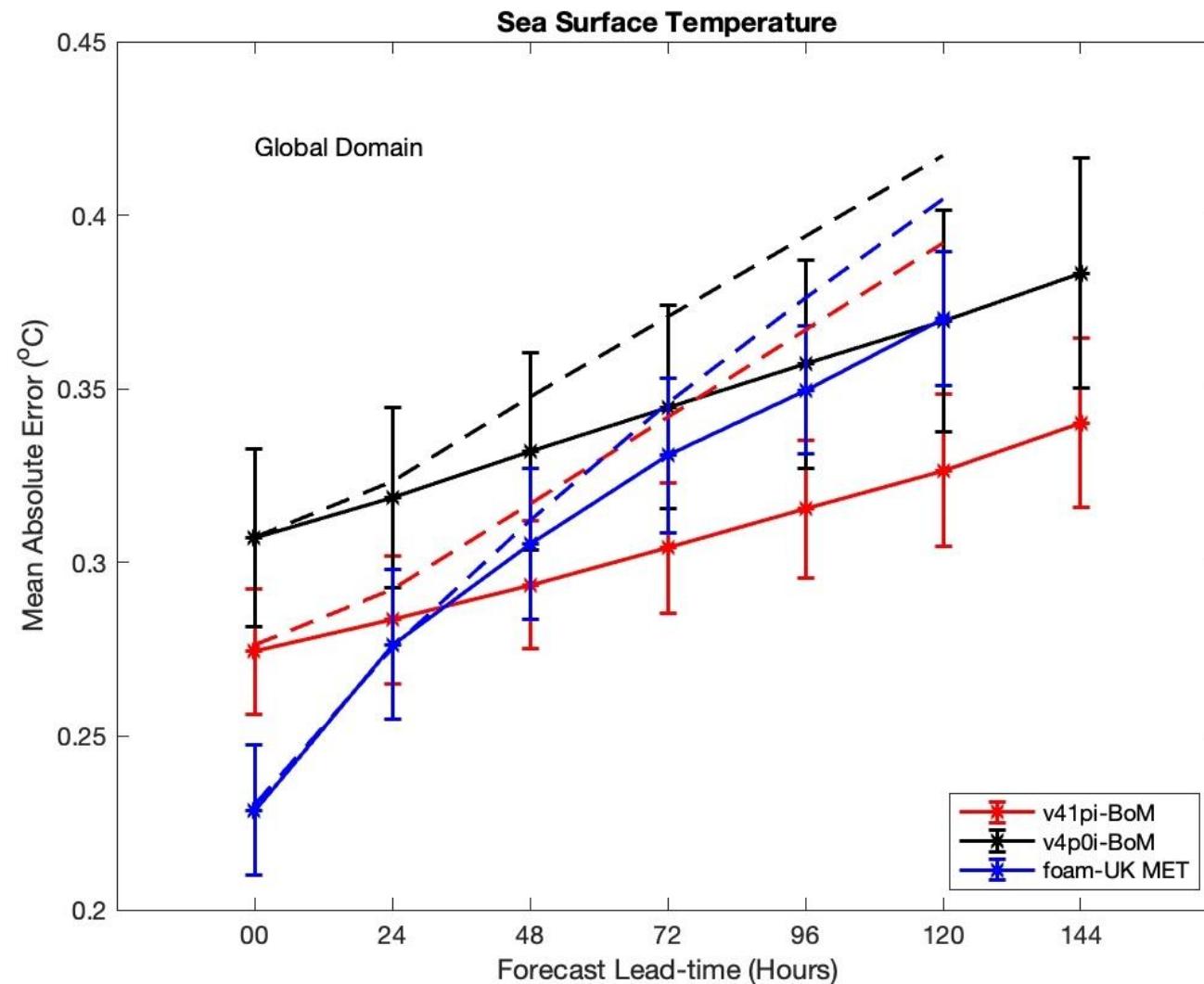


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OceanMAPSv4.1i



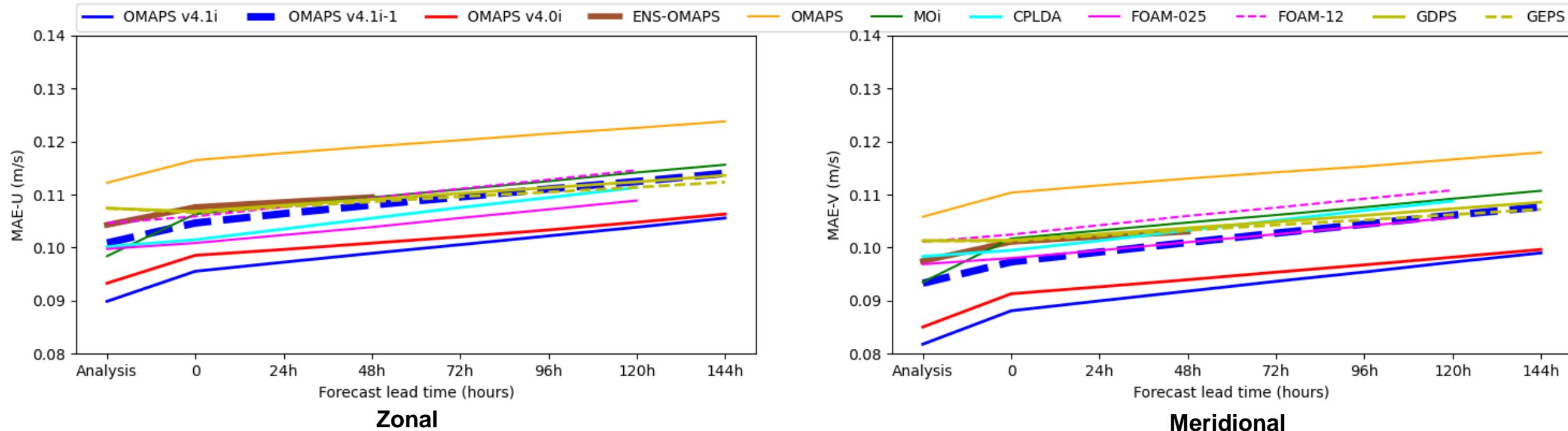
CLASS4 SST





Currents verification & intercomparison against global drifters (GDP)

Global forecast error growth





Conclusion

Transition to Hybrid-DLETKF was successfully implemented in 2022, (Brassington et al., 2023)
Updated to OceanMAPSv4.1i (1-day cycle)

Broad range of positive impacts

- Improved dynamical balance on increments
- Realistic separation of mesoscale eddies
- Adaptive to high variability/high tendency mesoscale dynamics
- Continuity of eddy propagation
- Constraining low signal to noise eddies
- Improved accuracy of state variables and ocean currents

NRT(-3)/BRT(-2) analysis design maximises observation coverage / minimise latency

Deterministic forecast from average restart initial conditions

- Improves forecast accuracy with minimal impact on mesoscale eddies





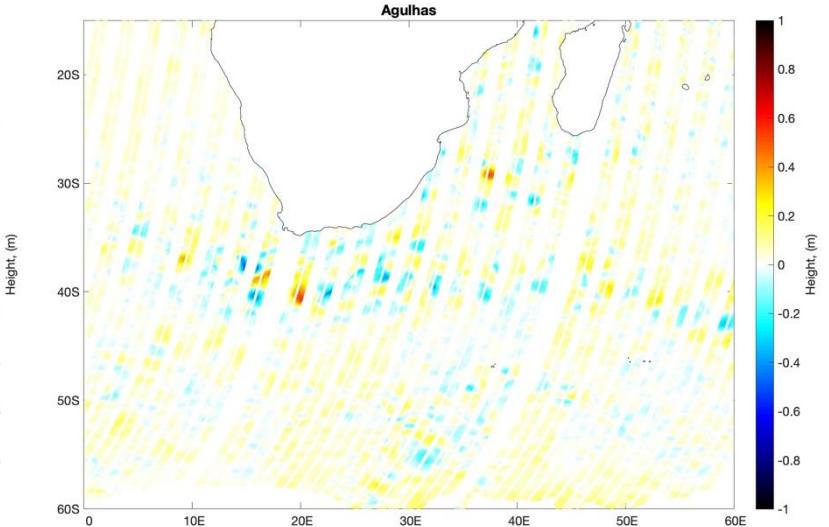
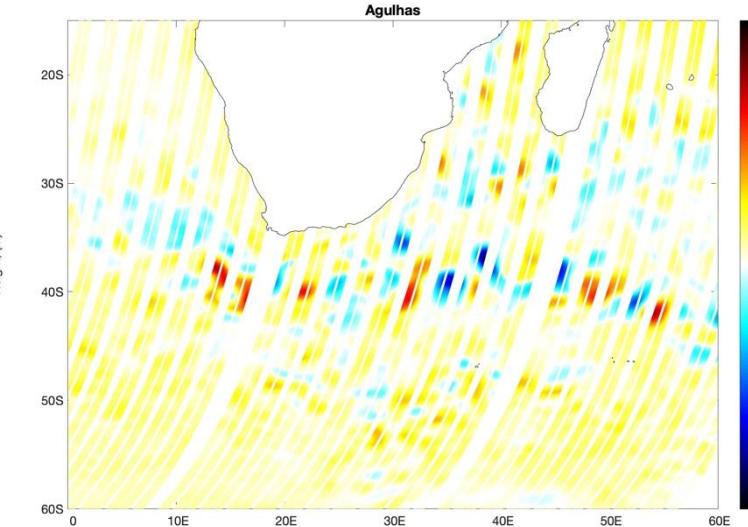
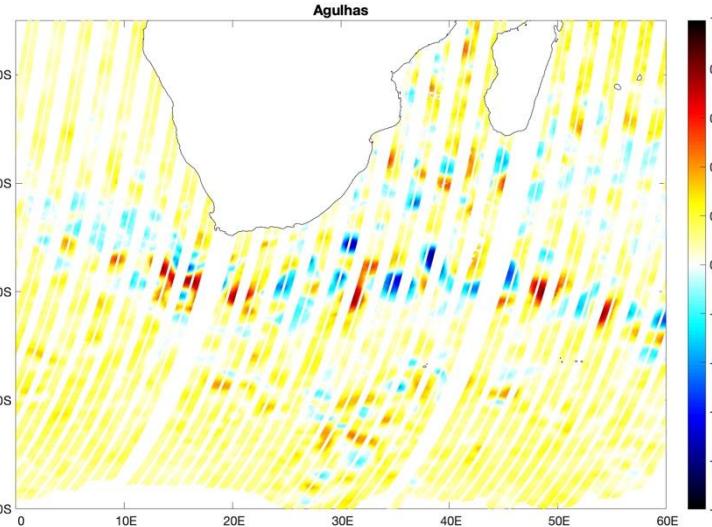
Thank you

Gary Brassington
gary.brassington@bom.gov.au

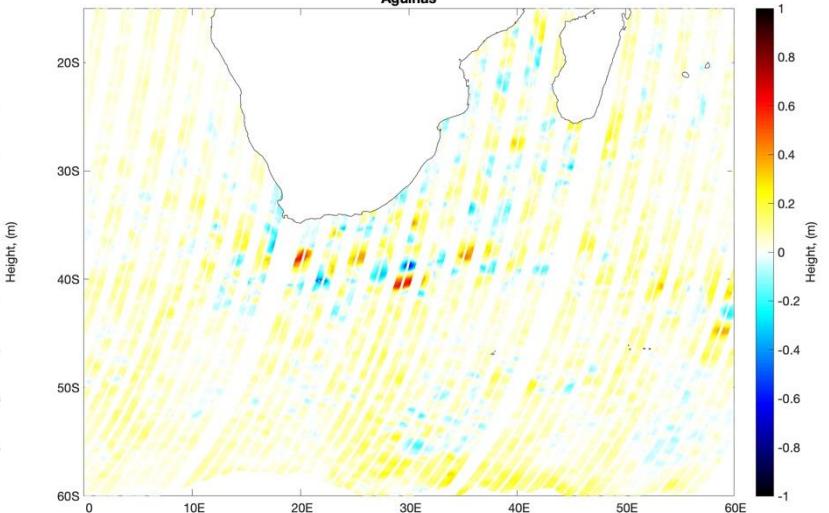
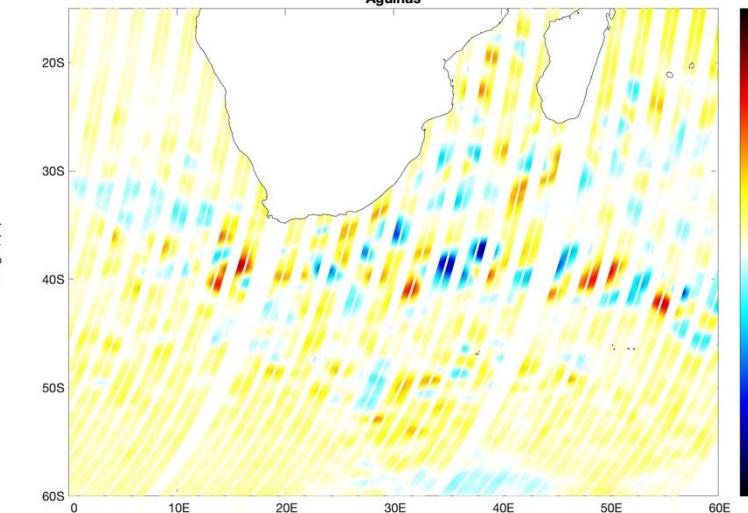
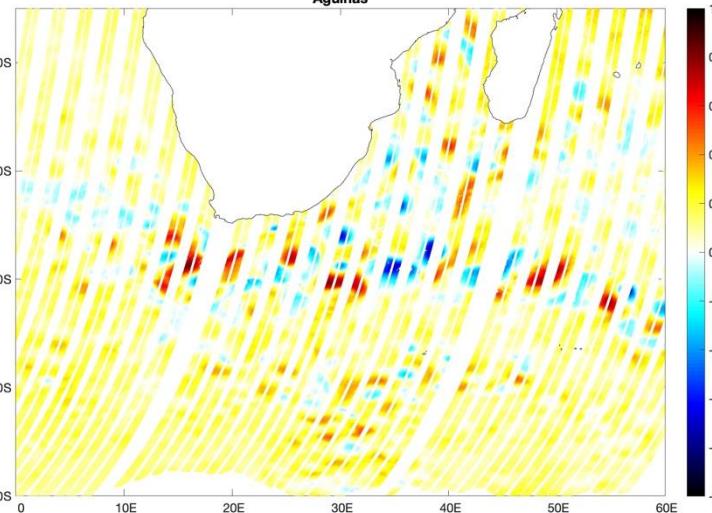
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Independent verification - SWOT

Cycle 005, Sub1, Asc.



Cycle 006, Sub1, Asc.



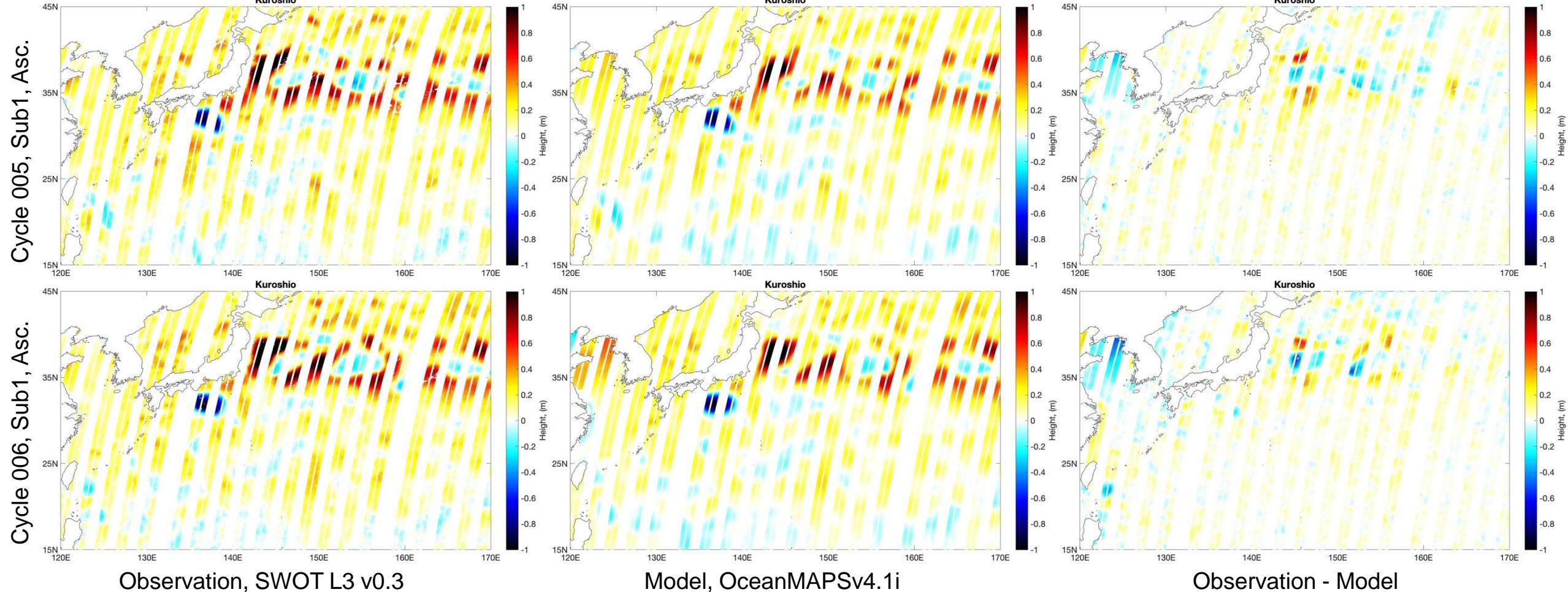
Observation, SWOT L3 v0.3

Model, OceanMAPSv4.1i

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Obs - Model

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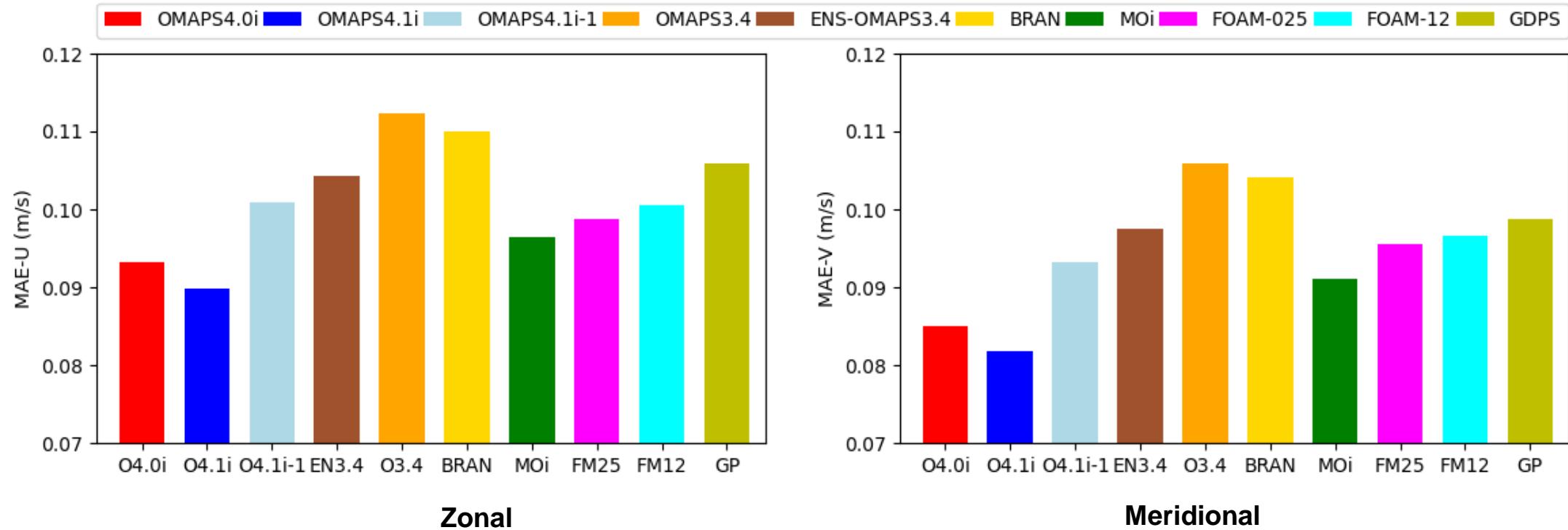


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Currents verification & intercomparison against global drifters (GDP)

Global Analysis



Zonal

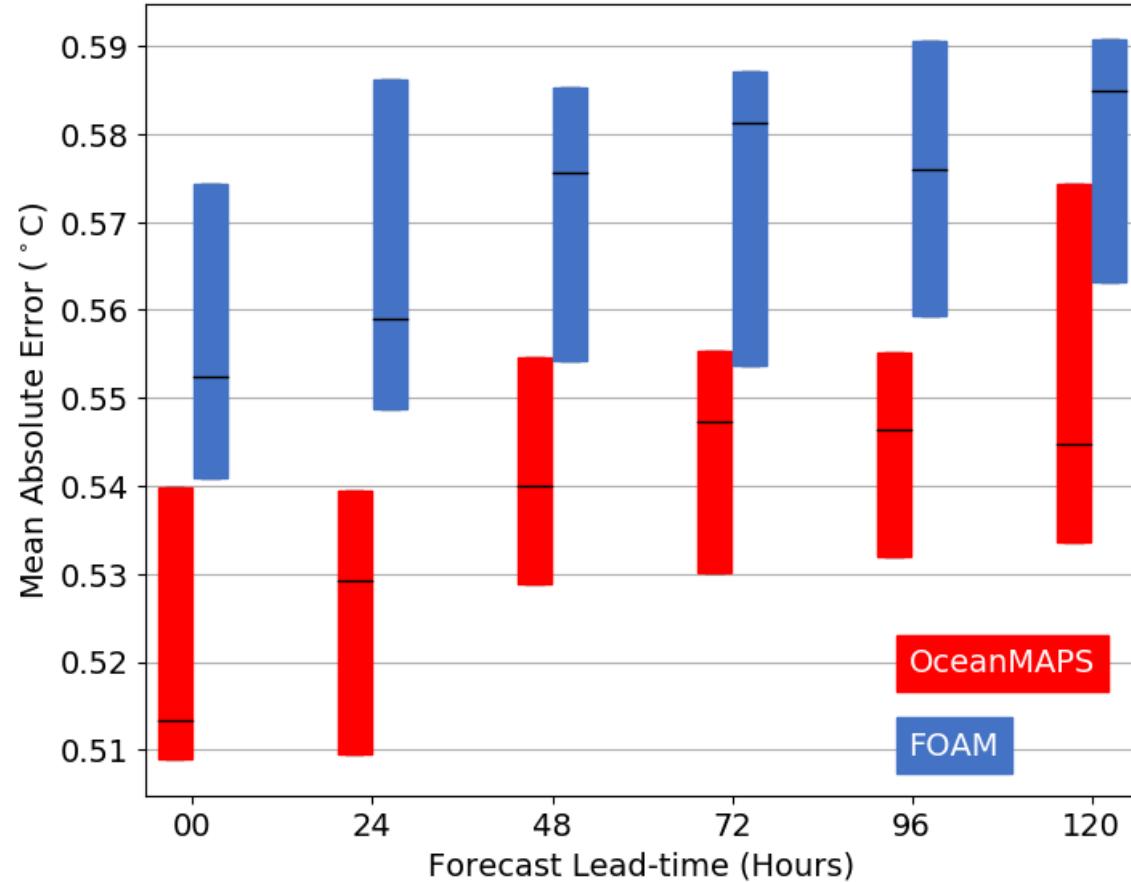
Meridional

OMAPS 4.0i, 4.1i, 4.1i-MEM1: 1JAN2024 – 30JUN2024 (182 days)
 MOi, FOAM-025, FOAM-12, GDPS: 1JAN2022 – 10MAR2023 (432 days)
 OMAPS3.4, ENS-OMAPS3.4: 20MAY2021 – 19MAY2022 (364 days)

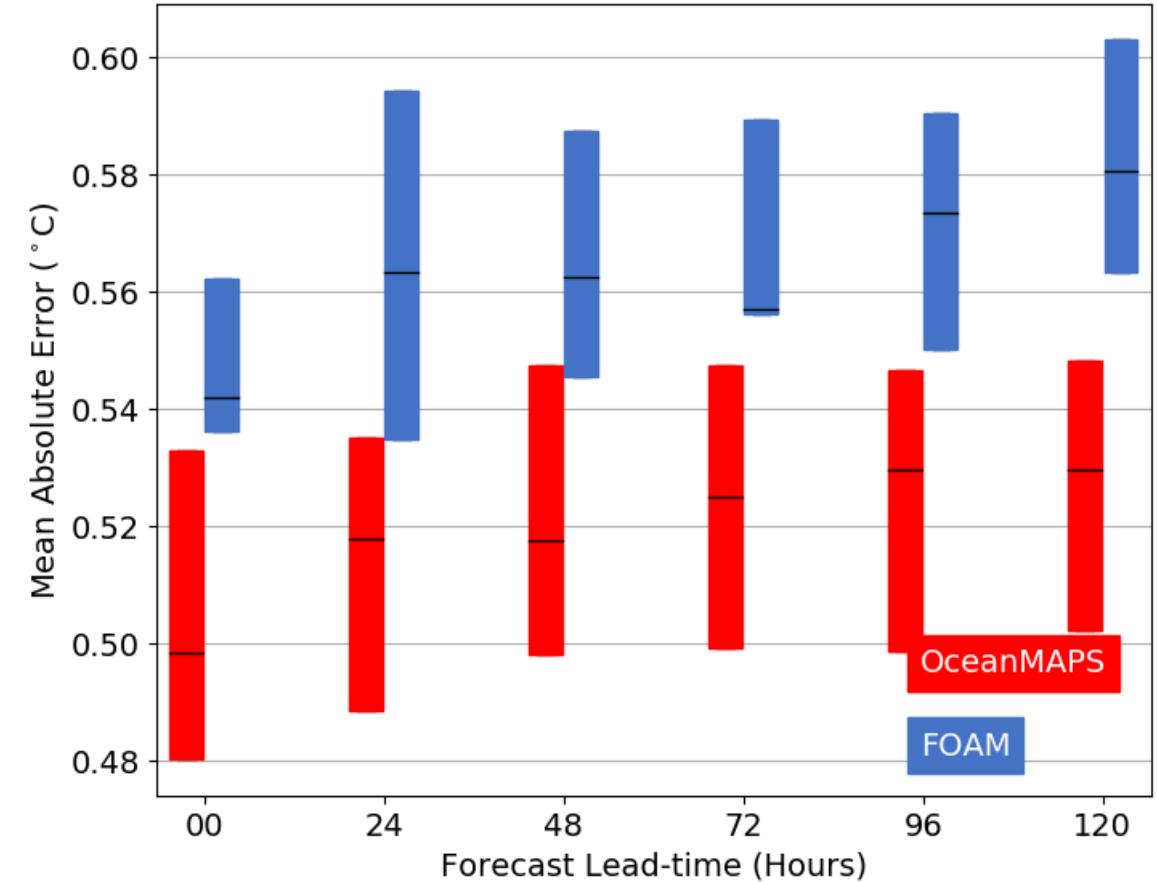
- Australia:** OMAPS3.4/4.0i/4.1i, ENS-OMAPS3.4, BRAN
France: MOi
UK: FOAM025, FOAM12
Canada: GDPS



Mean Absolute Errors, Global Temperature, March-April-May 2023



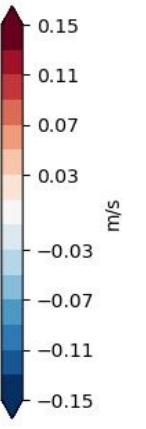
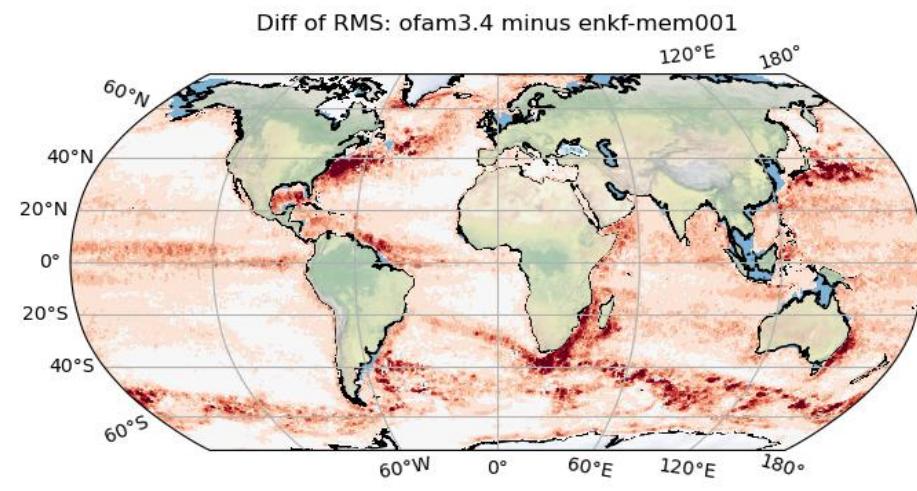
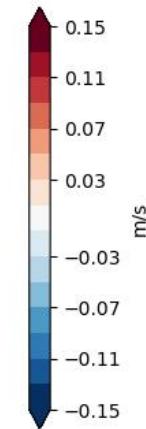
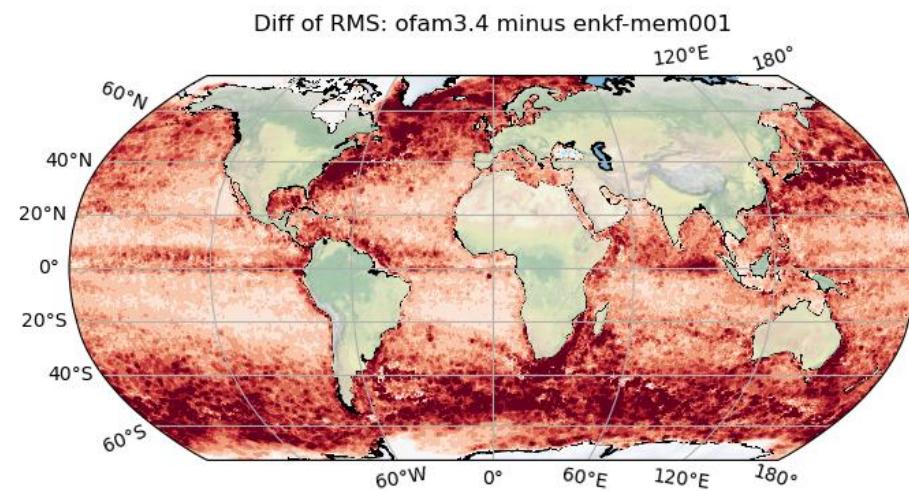
Mean Absolute Errors, Global Temperature, March-April-May 2024



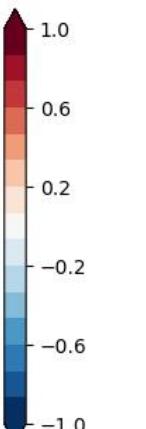
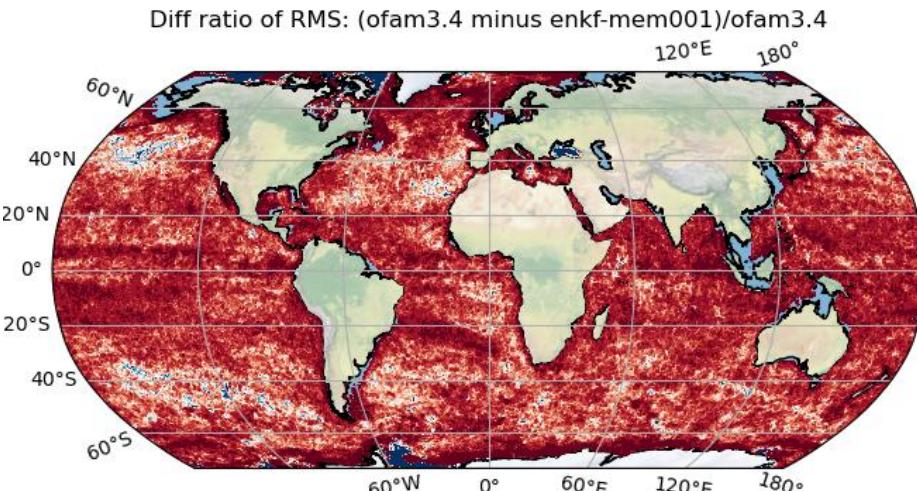
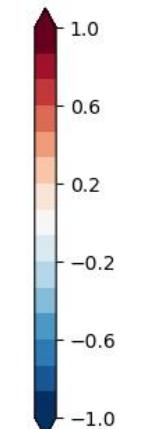
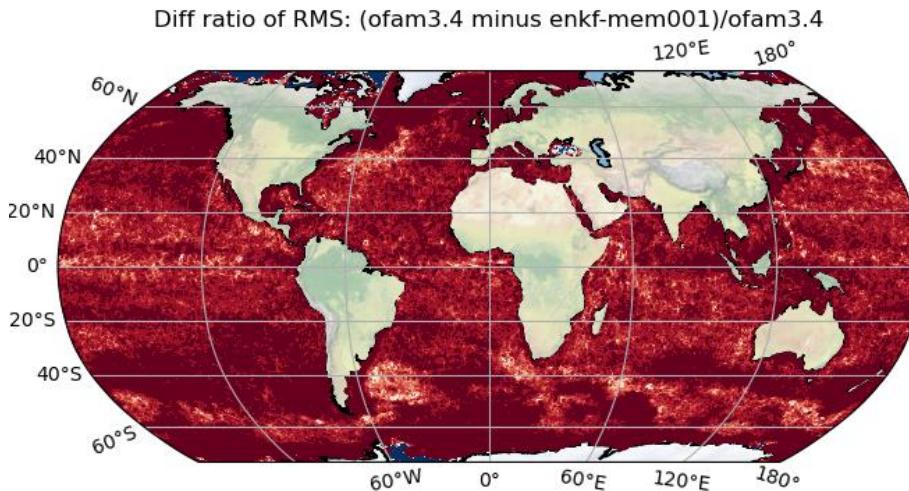
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$\Delta \text{RMS}(u_{\text{inc}})$



$\Delta \text{RMS} / \text{RMS}$



Surface

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95m

