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### **Ocean radiative transfer**

- Physical framework of ocean colour remote sensing
- Description of the radiance field in the ocean based on the Inherent Optical Properties (IOPs) of sea water
- We use a 1D 3-stream model coupled to a physical-BGC model to solve water columns individually. The aims are:
  - $_{\odot}$  To simulate radiometric variables that are comparable to observations
  - $_{\odot}$  To assess the influence of uncertainties in their simulation



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### Inputs

Direct and scattered (downward) surface irradiance



**IOPs** (absorption and scattering) We solve 4 major constituents:

- Pure water
- Phytoplankton Ο
- Non-algal particles Ο
- CDOM Ο



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### Outputs

#### In-water irradiance streams:

- Direct (downward)
- Scattered (downward)
- Backscattered (upward) -> Leaving-water irradiance

### Sea surface reflectance ( $R_{RS}$ )

- Ratio between upward and downward surface irradiance
- Model reflectance is transformed to remote-sensed reflectance by accounting for BRDF and the air-sea interface









### **Reflectance-derived chlorophyll**

We use a blue/green band ratio algorithm to derive reflectance-derived chlorophyll (rCHL), a proxy for surface chlorophyll.

$$log(rCHL) = \sum_{k=0}^{3} c_k \times \left[ log\left(\frac{R_{RS}(490)}{R_{RS}(555)}\right) \right]^k$$

with  $c_k$  coefficients as used in the CMEMS product for the Western Black Sea







### Uncertainties

- Here we consider uncertainties in the IOPs only, arising from:
  - Uncertainty on phytoplankton, non-algal particles and CDOM concentrations
  - $\circ$   $\,$  Uncertainty on the absorption and scattering spectra
- We use 1st order autoregressive processes following a log-normal distribution to perturb IOPs.
  - $\circ$  Time correlation: 30 days
  - $\circ$  Space correlation: 75 km
  - $\circ$  Standard deviation: 50%
  - o 20 members









## **Coupled modelling framework**



#### 1-way coupling:

- No feedback from the 3-stream RT model towards NEMO and the BGC
- RT model is used as an observation operator, projecting model variables into the space of observations





### **Optical contributions**



IOPs. 2017, eastern gyre

Absorption domination in 4 members: water, phytoplankton, non-algal particles, CDOM

2017-01-01



CDOM tends to dominate absorption outside of blooms, Phytoplankton and non-algal particles have IOPs of similar magnitude.









### Sea surface reflectance fields



2017 basin-wide average

Along BGC-Argo 6901866 track, 2017

- Overestimation of reflectance in longer wavelengths.
- rCHL is closer to measured chlorophyll during winter/spring blooms, but does not capture all observations.
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### rCHL distributions



- Initial bias in the deterministic solution of rCHL.
- In some conditions, distributions of rCHL are consistent with observations:
  - In the deep basin
  - In summer/autumn

- Observations are not properly captured in winter/spring. Reasons can be:
  - The deterministic solution is not good enough
  - $\circ$   $\,$  Other sources of uncertainty should be considered







# 2-way coupling ?

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- RT feedback to physics and BGC may improve the computation of R<sub>RS</sub> and rCHL.
- The influence on the physics and BGC has to be further assessed too.



## **Consequences of light perturbation**





Perturbation of surface irradiance - BGC-ARGO profile / Ensemble members

consequences on ocean biogeochemistry.

Perturbation of light (both directly and indirectly) can have major



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Conclusions



- The use of RT model shows the importance of modelling ٠ irradiance and IOPs (esp. CDOM) correctly.
- With rCHL, we obtain estimates of surface chlorophyll that are ۲ closer to observations that BAMHBI chlorophyll.
- Distributions of rCHL do not always capture observations, ٠ further calibration will be required for bloom periods.
- RT models allow to make a better use of reflectance products. ٠ Model calibration and validation
  - Data assimilation (poster by P. Verezemskaya, Monday)





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Thank you!







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